

# IGNEOUS PROVINCE ARE SIMILAR TO ISLAND ARC BASALTS?

Alexei V. Ivanov, Sergei V. Rasskazov, Elena I. Demonterova,  
Tatyana A. Yasnygina, Margarita N. Maslovskaya, Gennady D. Feoktistov  
Institute of the Earth's Crust, Siberian Branch, Russian Academy of Sciences

## STATEMENT OF THE PROBLEM

The Permo-Triassic Siberian Traps Large Igneous Province (STLIP), comprising volcanic and intrusive rocks from the Siberian Platform and West Siberian Basin (Fig. 1), is the most voluminous ( $>10^6 \text{ km}^3$ ) among known Phanerozoic large igneous provinces. Tholeiites and alkaline basalts with subordinate ultrabasic, alkaline, intermediate and acidic rocks make up the STLIP (Fig. 2a). According to  $\text{TiO}_2$  -  $\text{Mg\#}$  (where  $\text{Mg\#} = \text{Mg}/(\text{Mg} + 0.85\text{Fe}^{2+})$ ) diagram, tholeiites and alkali basalts can be subdivided into low-Ti and high-Ti series (Fig. 2b). Here we focus on rocks with  $\Delta \text{TiO}_2 < -0.4$  (where  $\Delta \text{TiO}_2$  is the deviation from the discrimination line in Fig. 3 calculated as  $\text{TiO}_2 = 3.45 - 0.0317 \times \text{Mg\#}$ ) (Fig. 2c). Such rocks are abundant in the STLIP. They are remarkably different from both the ocean-island-basalts (OIB) and enriched-middle-oceanic-arc-basalts (E-MORB). Instead they resemble island-arc-basalts (IAB) (Fig. 3), which origin is related with melting of water-rich upper mantle above subduction slabs. The question is why large portion of the intra continental STLIP basalts are IAB-like? Do they have any links with Permo-Triassic subduction processes? What is the mechanism of this link?

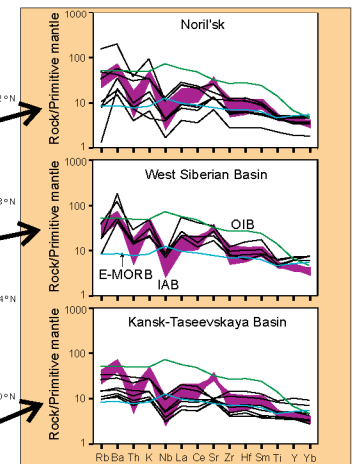
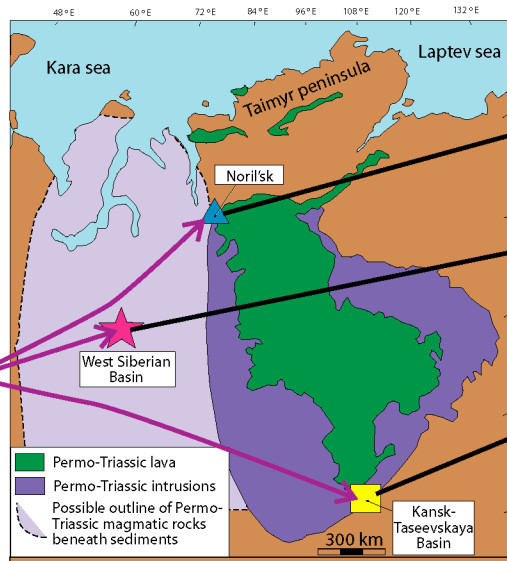
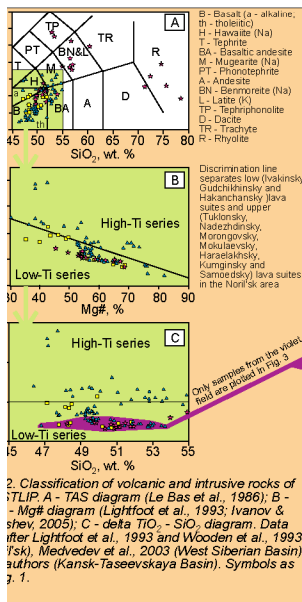
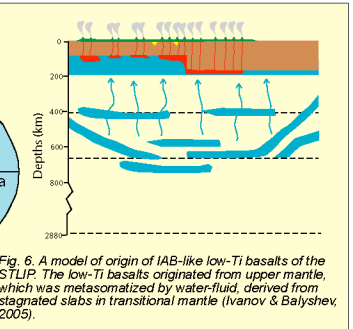
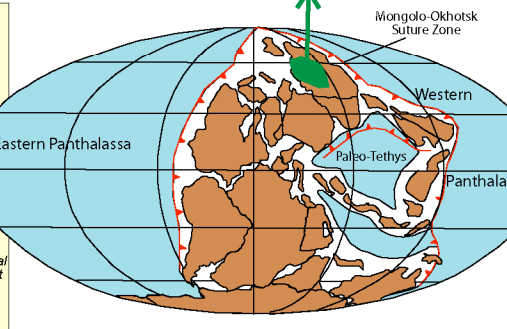
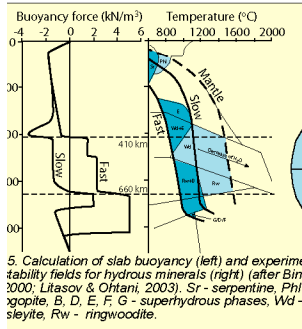


Fig. 1. Distribution of Late Permian and Early Triassic volcanic and intrusive rocks of the STLIP. Symbols mark three remote areas used for comparison of trace element geochemistry in low-Ti basalts (and dolerites).

Fig. 3. Primitive mantle trace element normalized diagram for low-Ti basalts (and dolerites) of the STLIP. Data are from Medvedev et al., 2003 (West Siberian Basin), Lightfoot et al., 1993 (Noril'sk) and authors (Kansk-Taseevskaya Basin). OIB and E-MORB are average compositions after Sun and McDonough, 1989. IAB is the field of the subduction-related basalts of the Quaternary volcanic belt of the Sredinny Ridge, Kamchatka (Churikova et al., 2001; Ivanov et al., 2004).



5. Calculation of slab buoyancy (left) and experimental stability fields for hydrous minerals (right) (after Bina et al., 2000; Litasov & Ohtani, 2003). Sp - serpentine, Phl - phlogopite, Sph - sphalerite, Ring - ringwoodite.

Fig. 4. Tectonic reconstruction map (redrawn from Nikishin et al., 2002). STLIP is shown in green.

Fig. 6. A model of origin of IAB-like low-Ti basalts of the STLIP. The low-Ti basalts originated from upper mantle, which was metasomatized by water-fluid, derived from stagnated slabs in transitional mantle (Ivanov & Balychev, 2005).

## DISCUSSION

Permian, Siberian part of Pangaea was surrounded by subducting systems (Fig. 4). We suggest that subducting slabs brought water into transitional zone of the mantle. Experimental data show that superhydrous minerals are stable at fast subducting slab conditions within two mantle regions at about 410 and 660 km discontinuities (Fig. 5 right). Fast subducting slabs can not penetrate through the 660 km discontinuity, because they attain positive buoyancy (Fig. 5 left). Therefore, slabs that stagnated in the transitional mantle will enrich the mantle in water. With heating slabs up the water will be released due to decomposition of the superhydrous minerals, followed by water saturation of the upper mantle and consequent voluminous melting (Fig. 6).

## REFERENCES

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